

A Coruña, 14 diciembre 2010

"El futuro del clima: proyecciones, incertidumbres y cambios abruptos"

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GOBIERNO
DE ESPAÑA

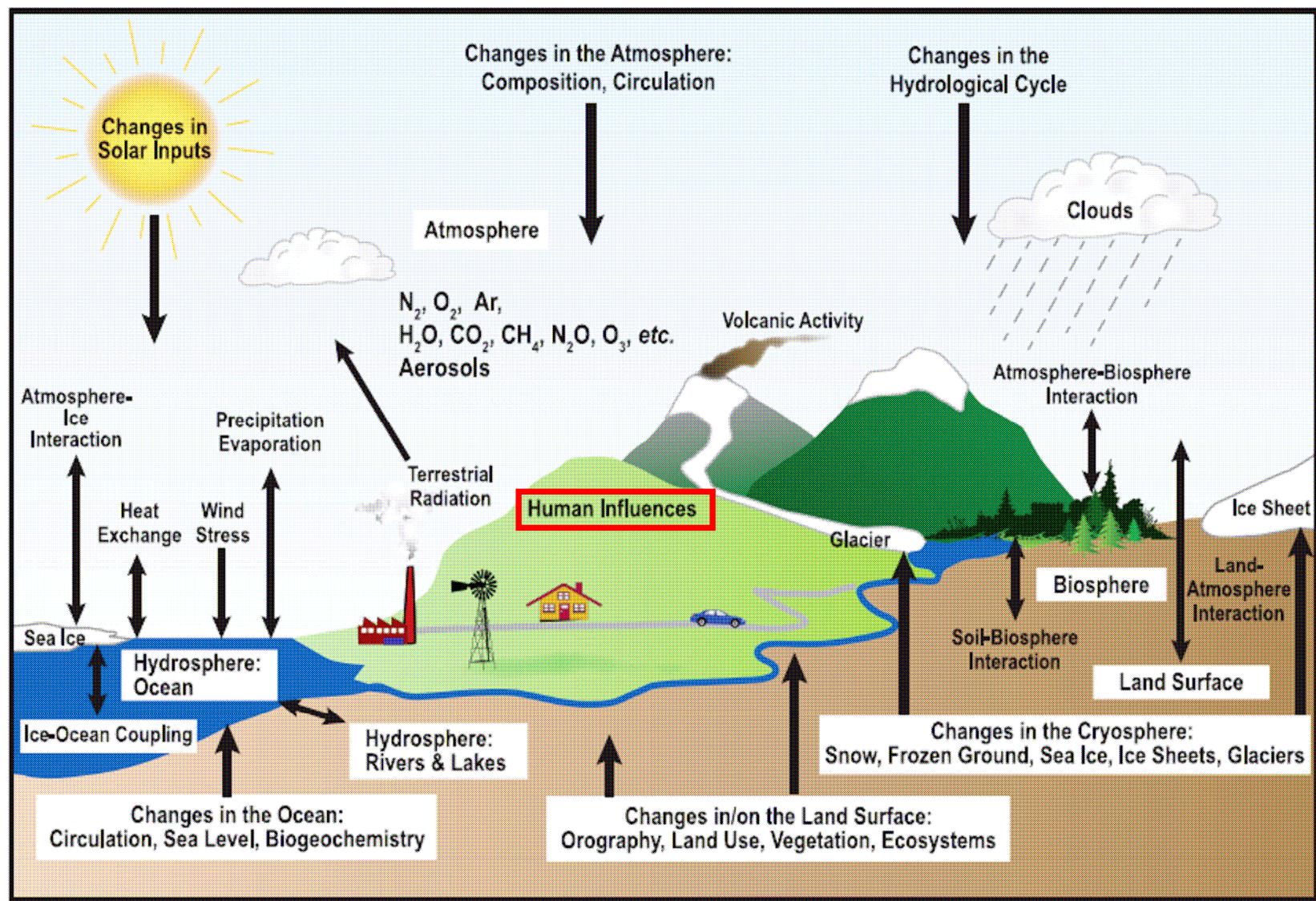
MINISTERIO
DE MEDIO AMBIENTE
Y MEDIO RURAL Y MARINO

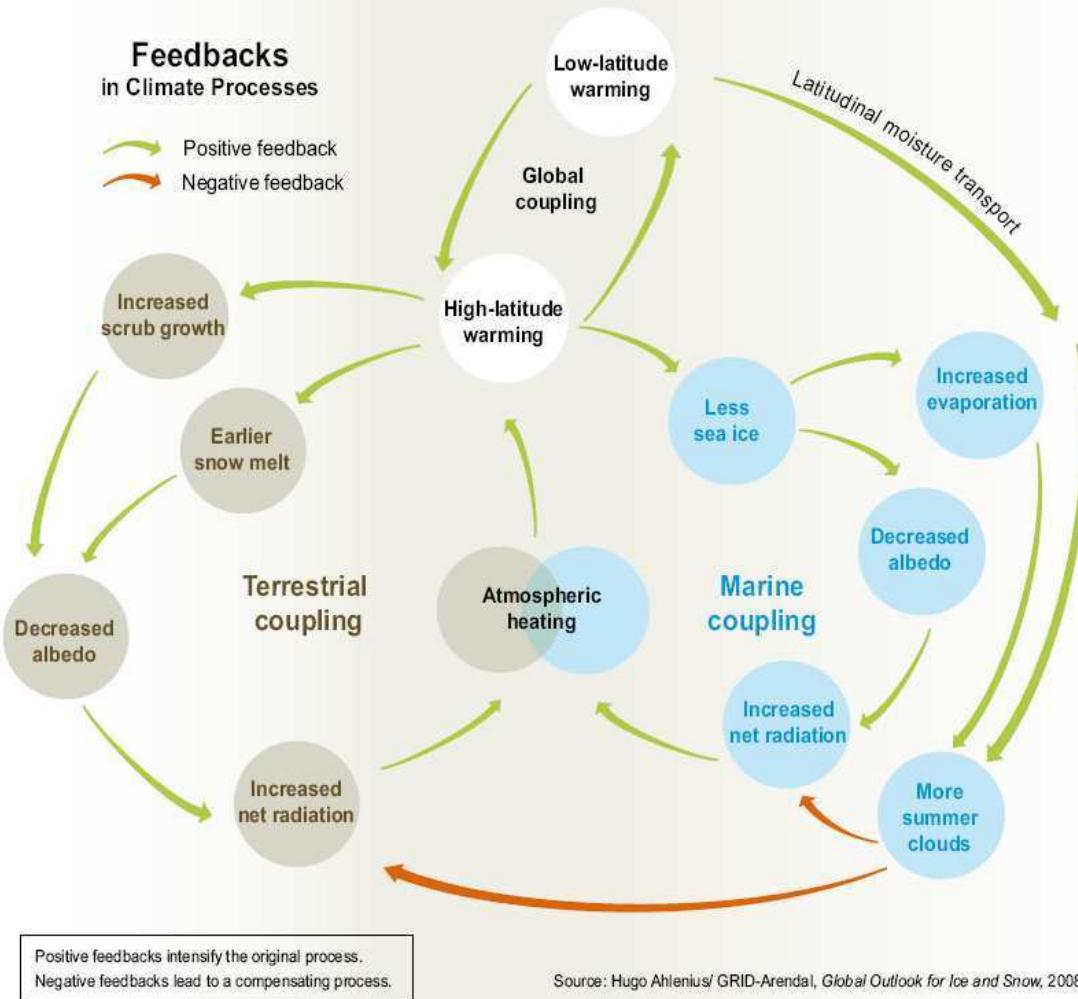


Contenido

- Introducción
- Observación del cambio climático
- Proyecciones climáticas
- Causas
- Incertidumbres
- Cambios abruptos
- Futuro

Sistema climático





- Climate system is highly nonlinear
- Strong coupling among subsystems with different time scales

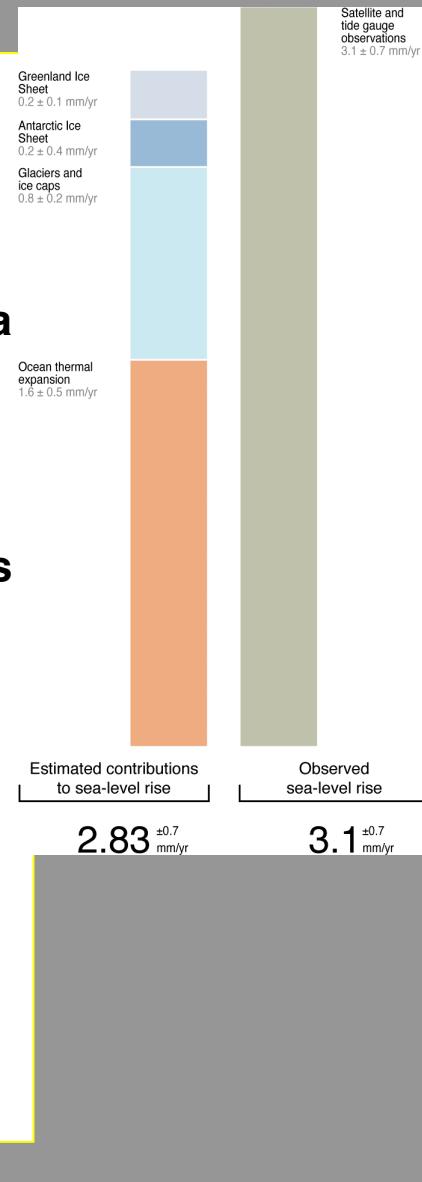
Models needed!

¿Qué CAMBIO CLIMATICO se observa?

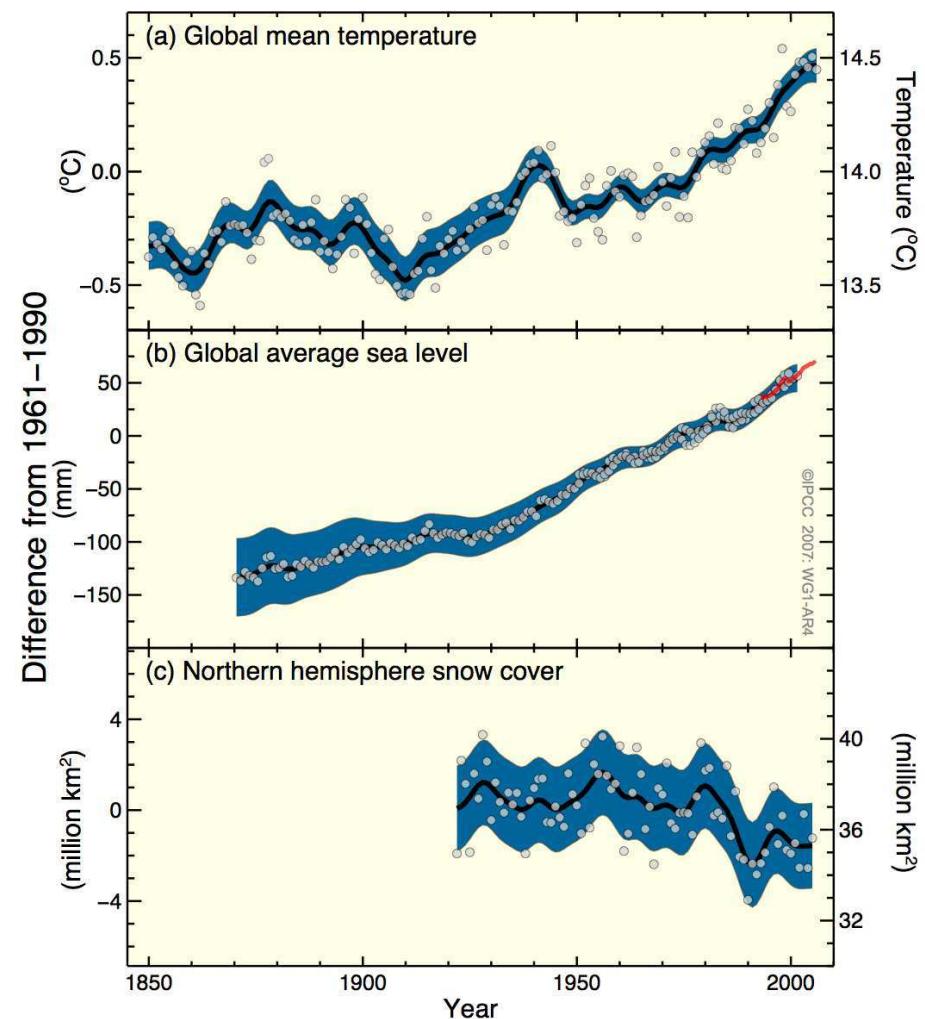
El **calentamiento del sistema climático es inequívoco**,

tal y como se evidencia de las observaciones del incremento de las temperaturas globales medias del aire y del océano, de la fusión de las nieves y hielos y de la elevación global del nivel medio del mar

(IPCC-AR4 *dixit*)

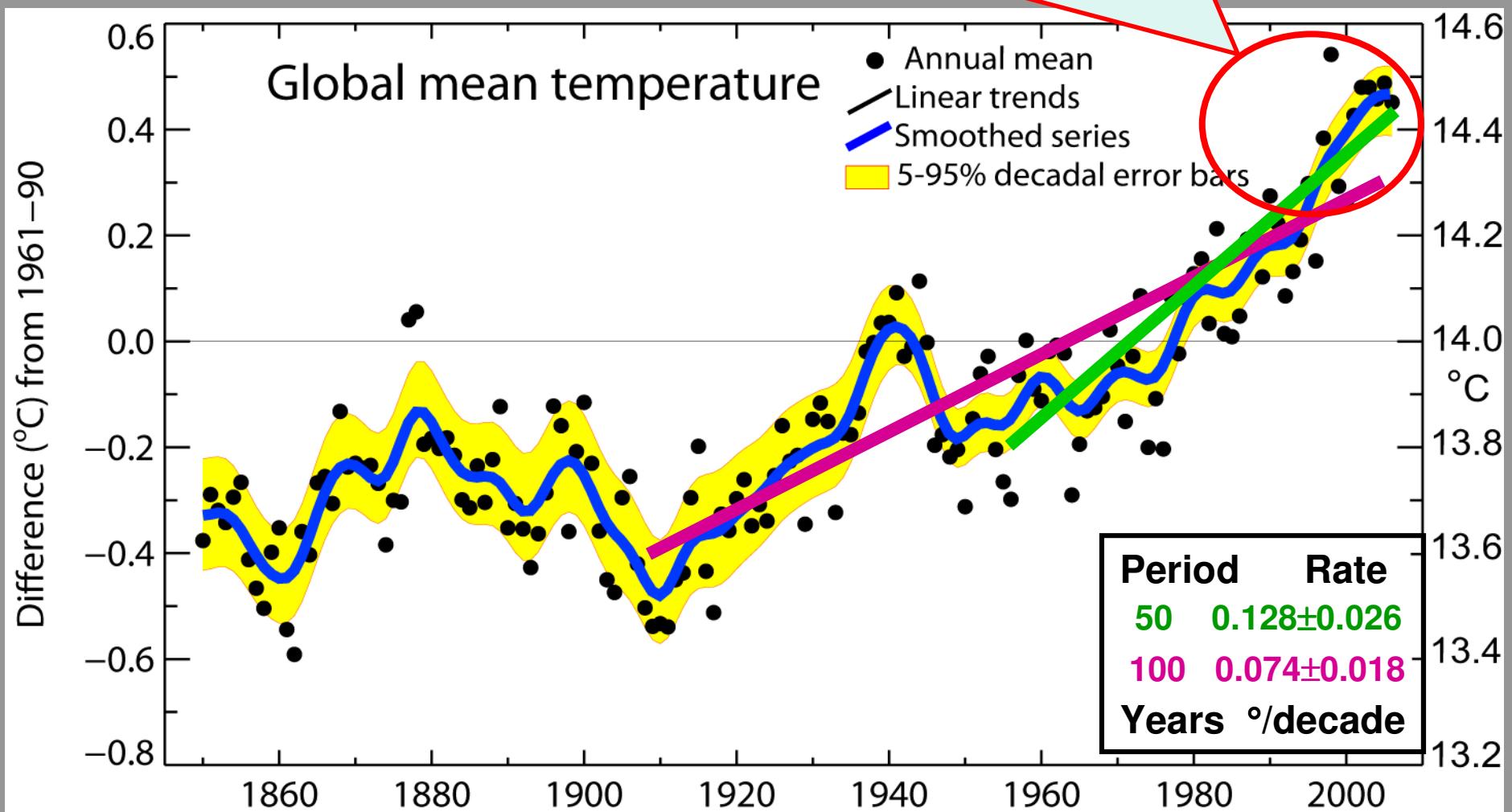


Changes in Temperature , Sea Level and Northern Hemisphere Snow Cover



Temperaturas globales

Los 12 años más cálidos:
1998, 2005, 2003, 2002, 2004, 2006,
2001, 1997, 1995, 1999, 1990, 2000



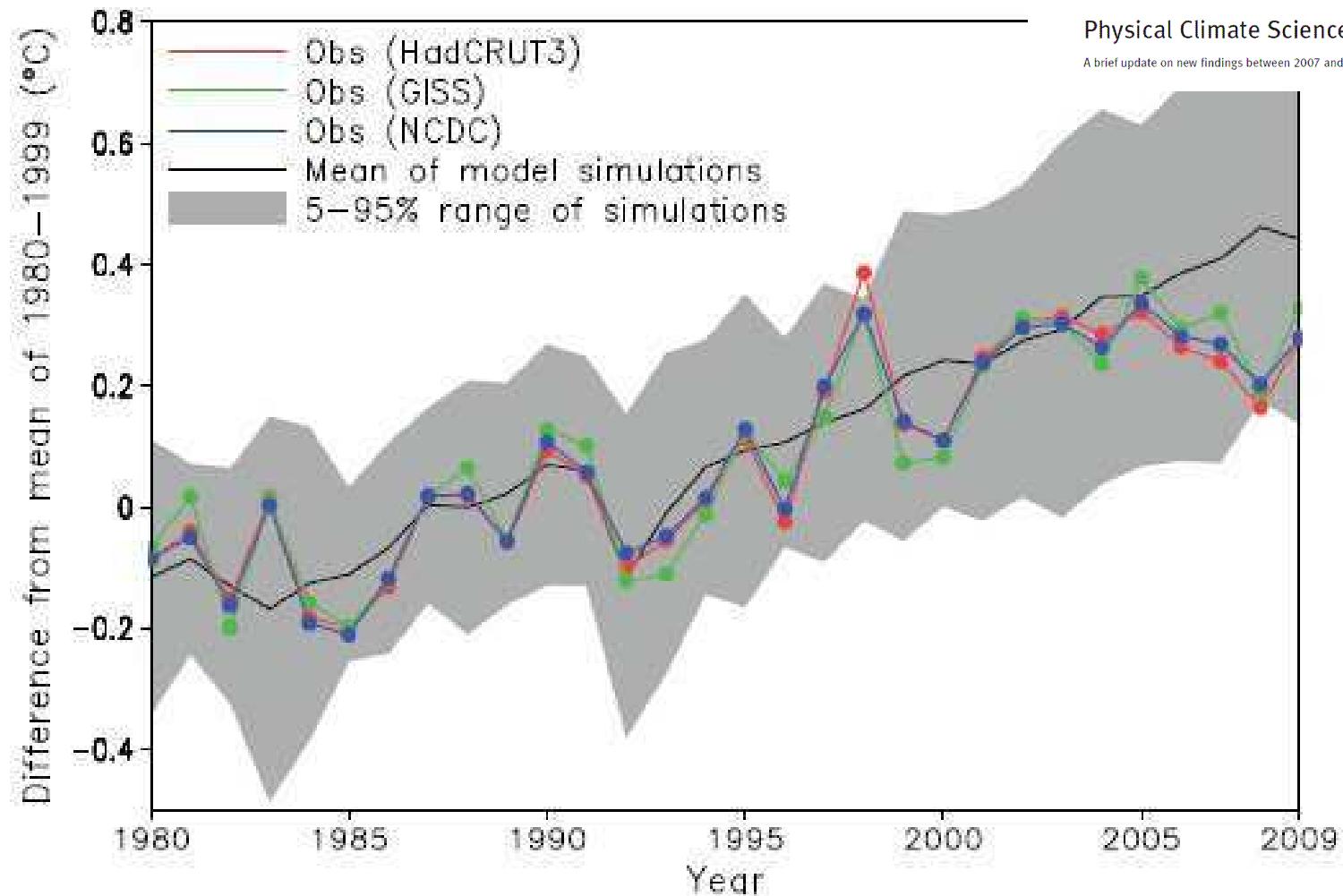
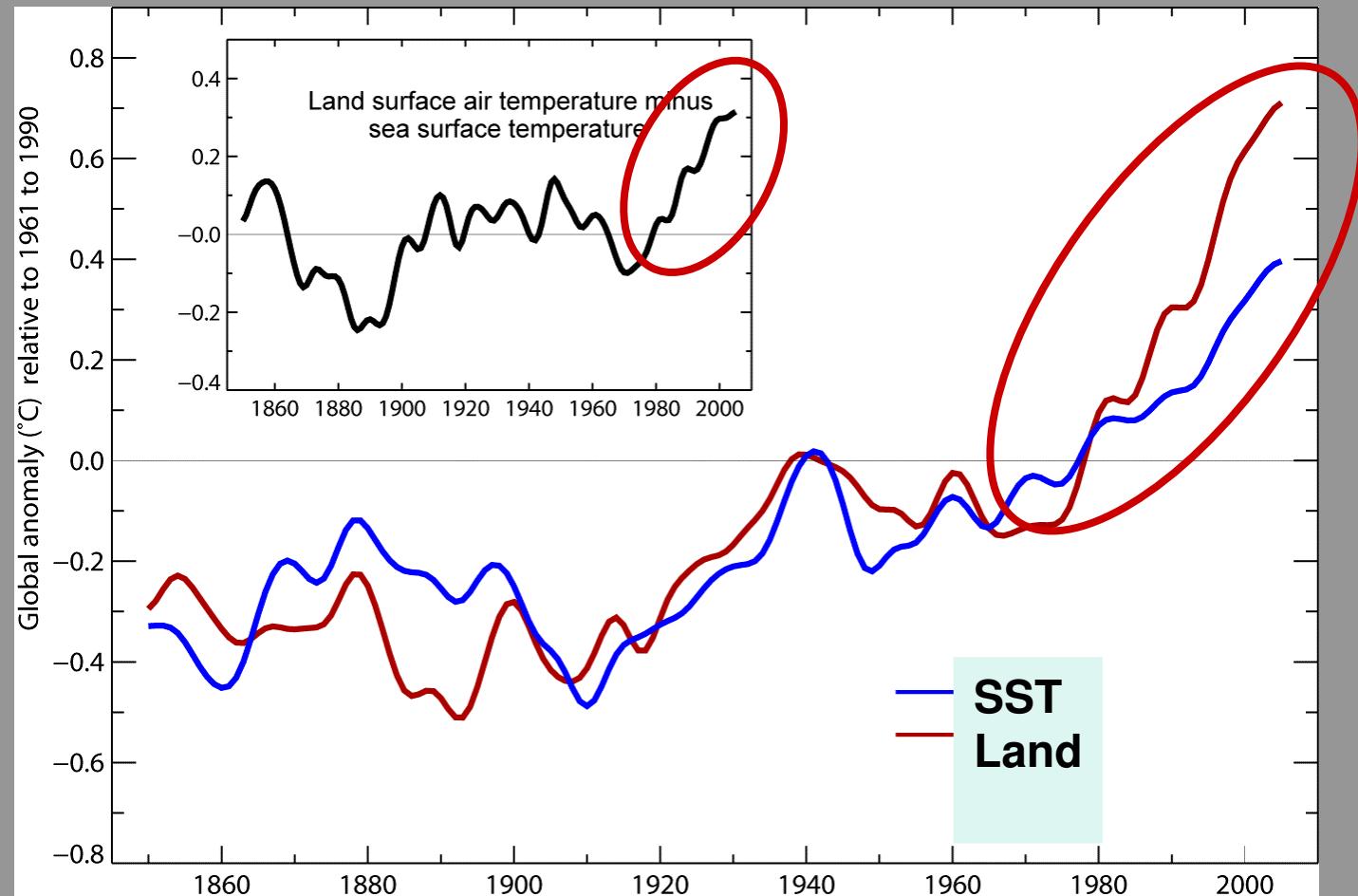
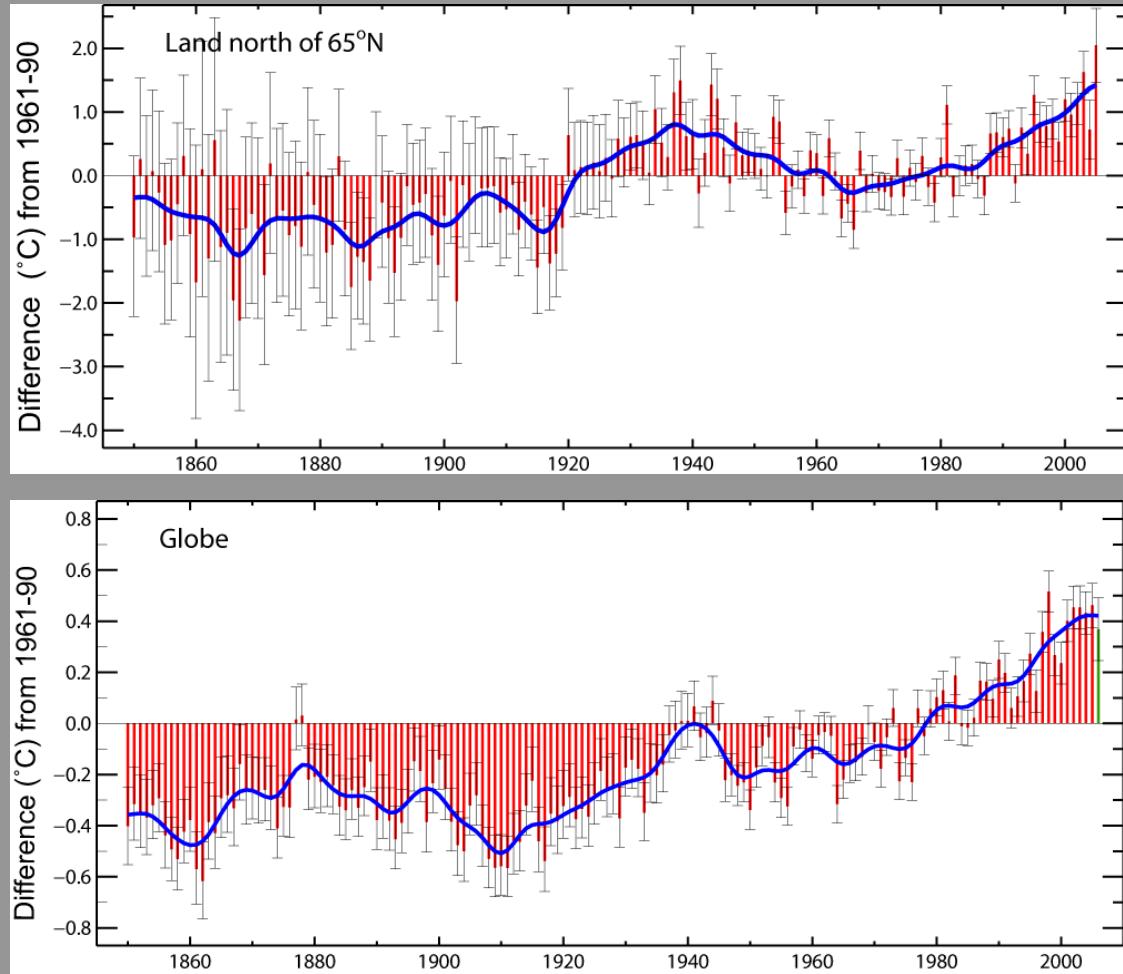


Figure 2. Recent changes in global mean temperature according to three analyses of observations (HadCRUT3, GISS and NCDC) and the 21 climate model simulations without climate policies used in IPCC AR4¹⁹. The mean of the model results is given by the solid black line. The shading indicates a 5–95% range of variability according to the same model simulations. The temperatures are given as deviations from the mean value over the period 1980–1999.

Temperatura superficie terrestre suben más deprisa que SST



Anomalías de temperatura anual Ártica vs Global (°C)

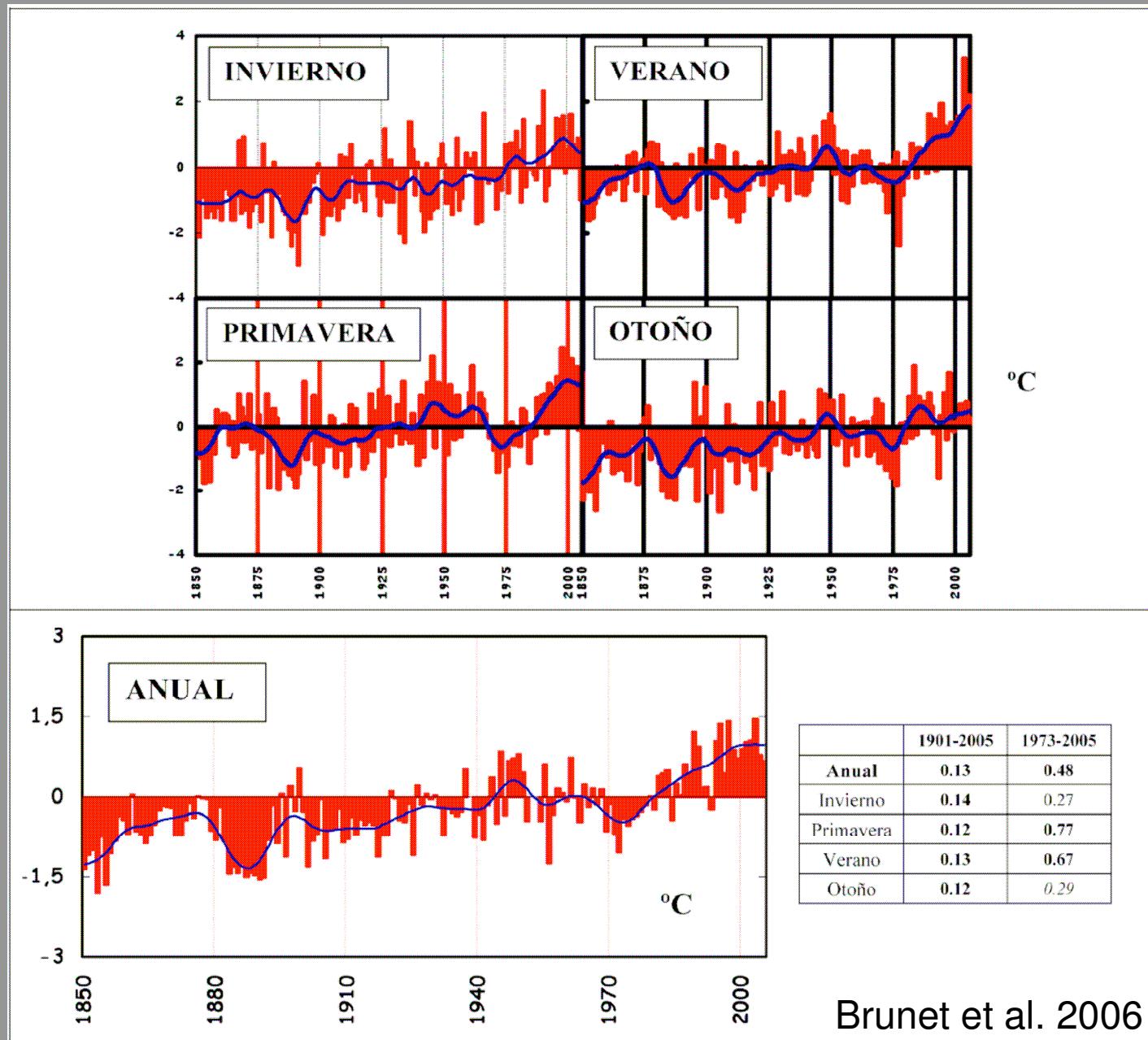


Calentamiento en el Ártico es doble que para el globo desde el s. XIX al XXI y desde los últimos 1960s hasta la actualidad.

Calentamiento entre 1925 y 1950 en el Ártico no fue tan generalizado como el reciente calentamiento global.

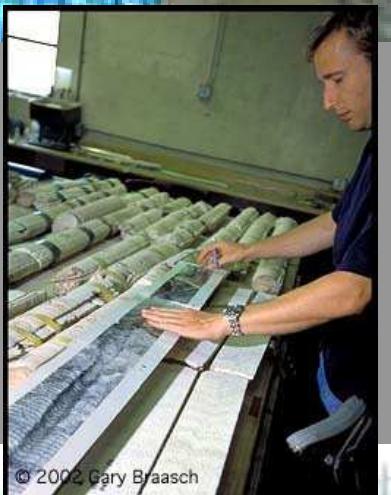
Escalas diferentes!

Cambio de temperaturas en España (I)





OBSERVACION DEL CLIMA (II)



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TABLE 2. Major Paleoclimatic Data Sources

Data sources	Variable measured	Continuity of evidence	Potential geographical coverage	Period open to study (years)	Minimum sampling interval (years)	Usual dating accuracy (years)	Climate-related inferences
Ocean sediments	Isotopic composition of planktonic and benthic fossils Floral and faunal assemblages Morphological characteristics of fossils Mineralogical composition and abundance	Continuous	Global ocean, except (for carbonate fossils) deepest zones (below CaCO ₃ compensation depths)				Global ice volume; surface temperature and salinity; bottom temperature and bottom water flux; aridity of adjacent land areas; prevailing wind direction and strength
	Sedimentation rates (cm per 1000 years)						
	<2	Favored areas along continental margins	100 000+	1000+	±5%	±5%	
	2-5		200 000+	500+	±5%	±5%	
	>10		10 000+	50+	±5%	±5%	
Ice cores	Oxygen isotope composition Trace chemistry and electrolytic conductivity Fabric	Continuous	Glaciated regions in polar and alpine areas (optimally in dry snow zones)	100 000+	Variable, but optimally 1-10 years for last 10 ⁴ years	Variable, but optimally 0.05% for last 10 ⁴ years	Temperature, accumulation rates atmospheric composition and turbidity, ice thickness (height), solar output variations
Mountain glaciers	Terminal positions Glaciation levels and equilibrium line altitudes	Episodic	45°S to 70°N	50 000	—	±5-10%	Temperature, precipitation (net accumulation)
Closed basin lakes	Lake level	Episodic	Low to mid latitudes (arid and semi-arid environments)	50 000	—	±5%	Moisture availability ("effective precipitation")
Bog or lake sediments (varved sediments)	Insect assemblage composition Pollen type concentration, geochemical and sedimentological composition	Continuous	All continents Mid to high latitudes	10 000+ (common) 150 000 (rare)	~50 1-10	±5% +1-10	Temperature, precipitation, soil moisture, air mass frequencies
Tree rings	Ring width anomaly, density, isotopic composition	Continuous	Mid- and high-latitude continents	1000 (common) 8000 (rare)	1	1	Temperature, precipitation, moisture, (circulation)
Written records	Phenology, weather logs, sailing logs, etc.	Episodic or continuous	Global	1000+	1	1	Varied

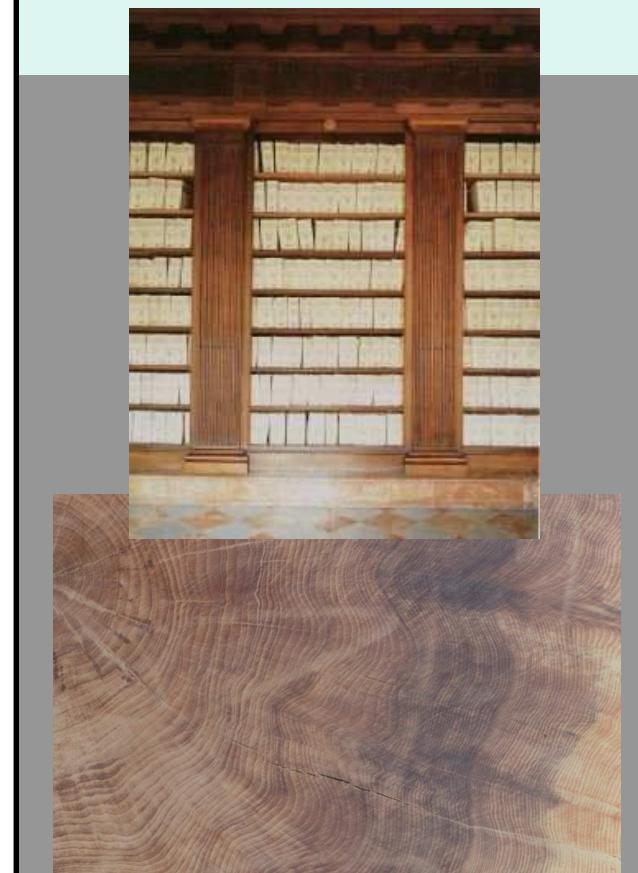
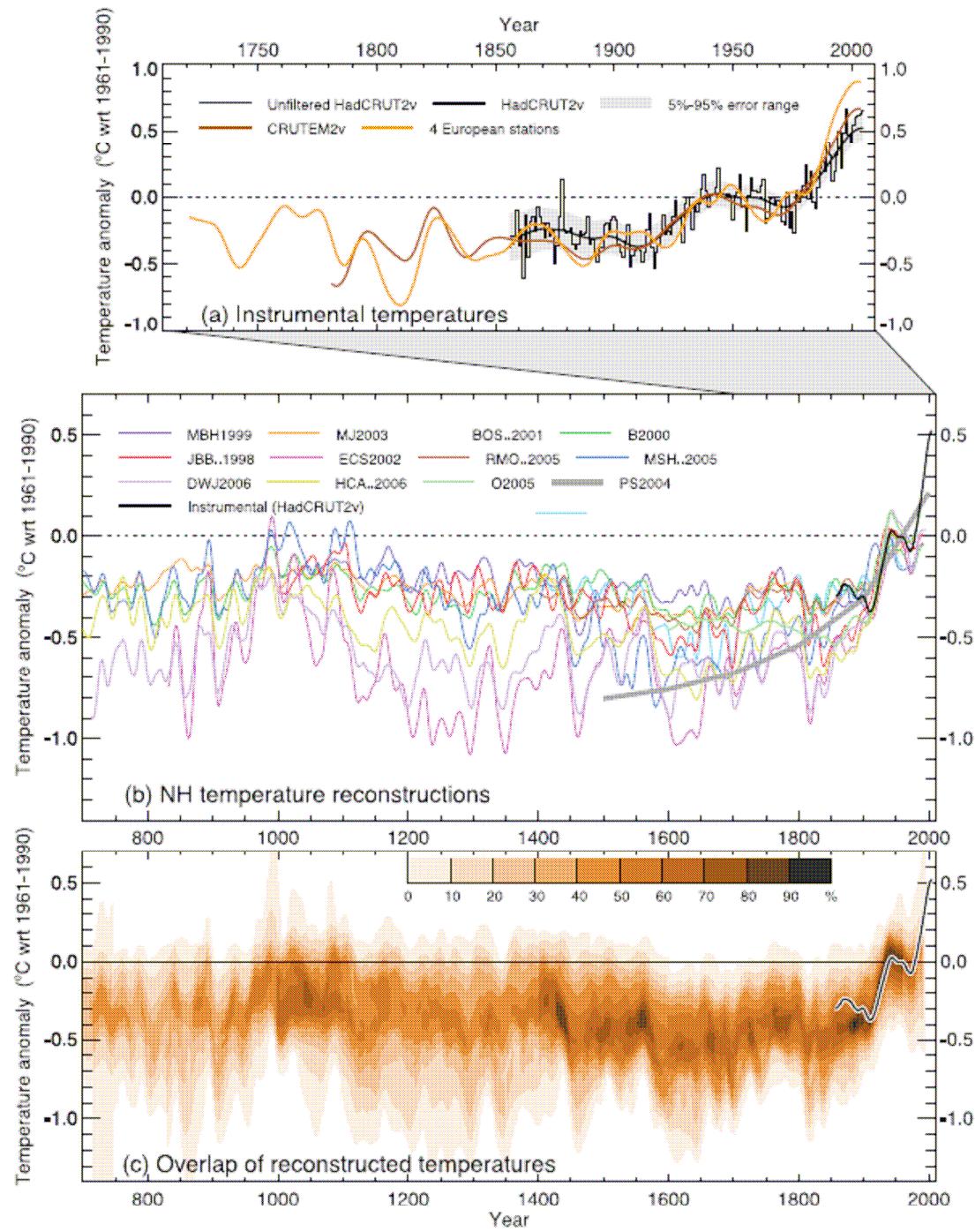


<http://www.ucm.es/info/reclido/>



<http://www.worldviewofglobalwarming.org/index.html>

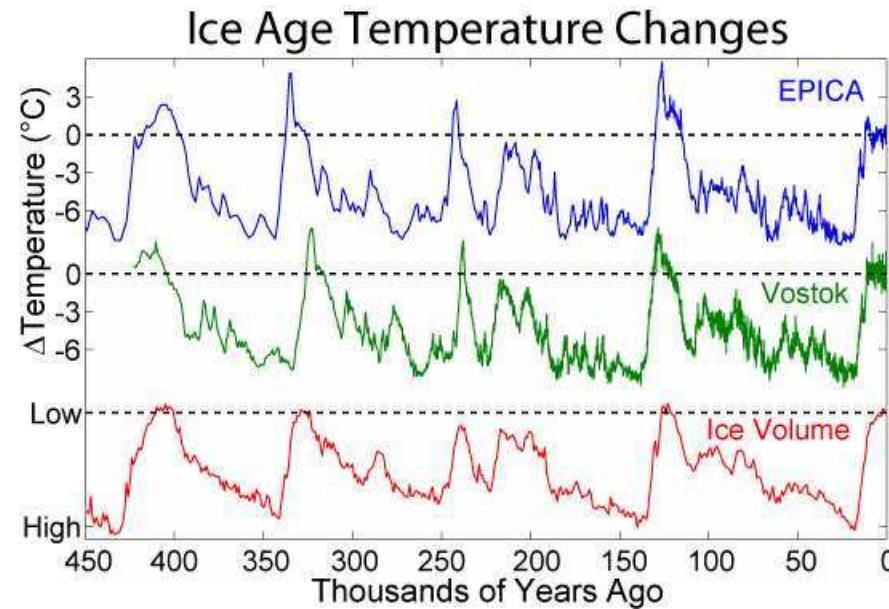
Los últimos 2000 años (IPCC-AR4, 2007)



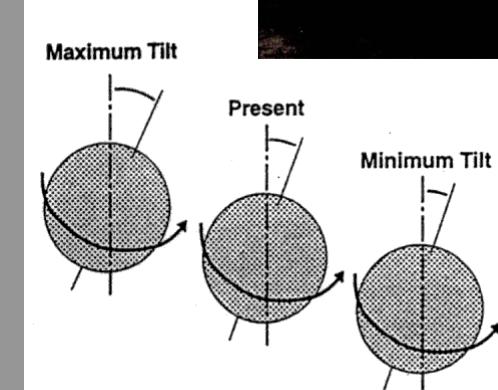
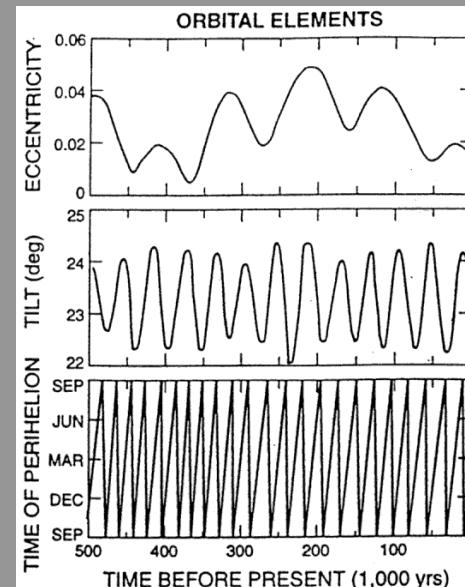
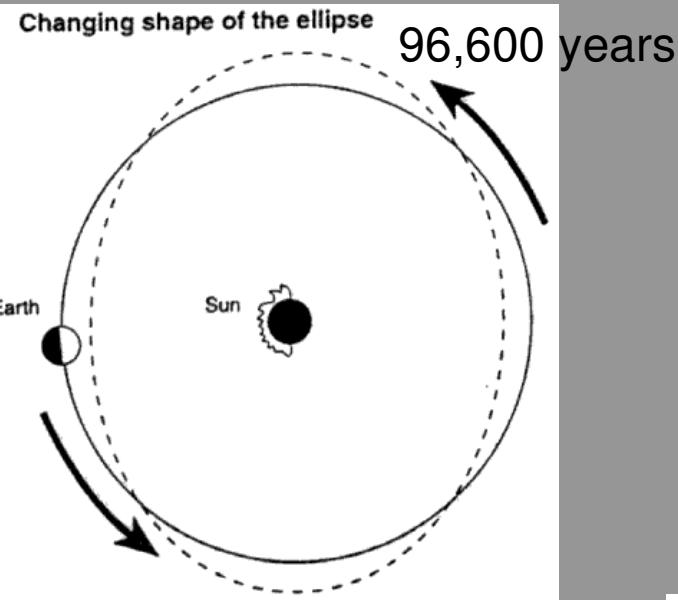
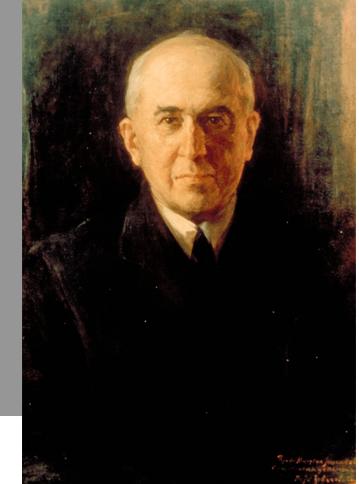


Louis
Agassiz
(1837)

- He was the first to scientifically propose that the Earth had been subject to a past ice age.



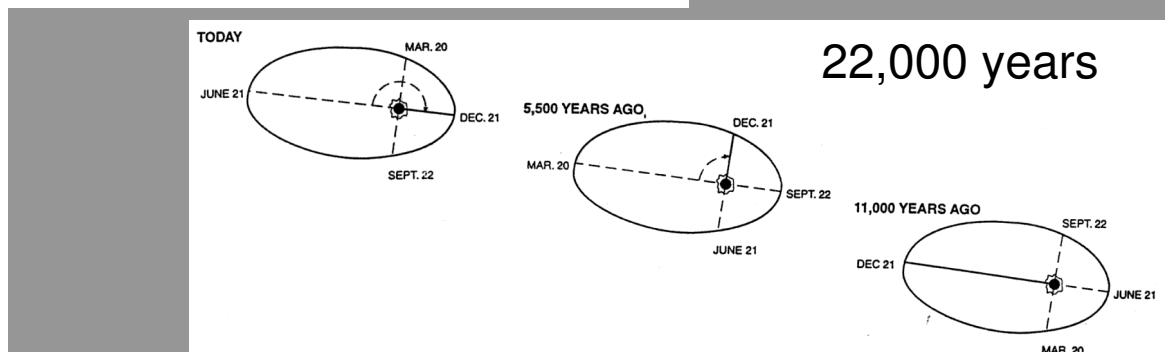
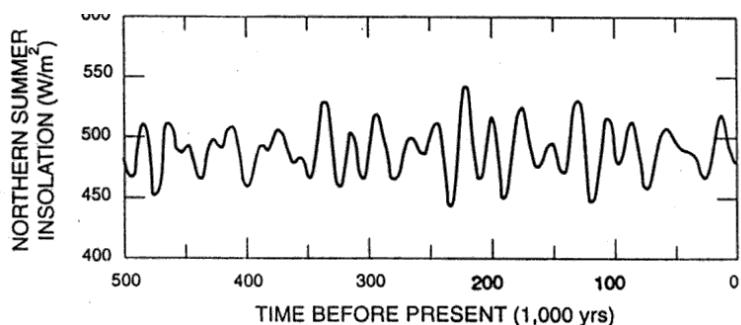
FORZAMIENTO NATURAL DEL SISTEMA CLIMATICO: ciclos de Milankovitch



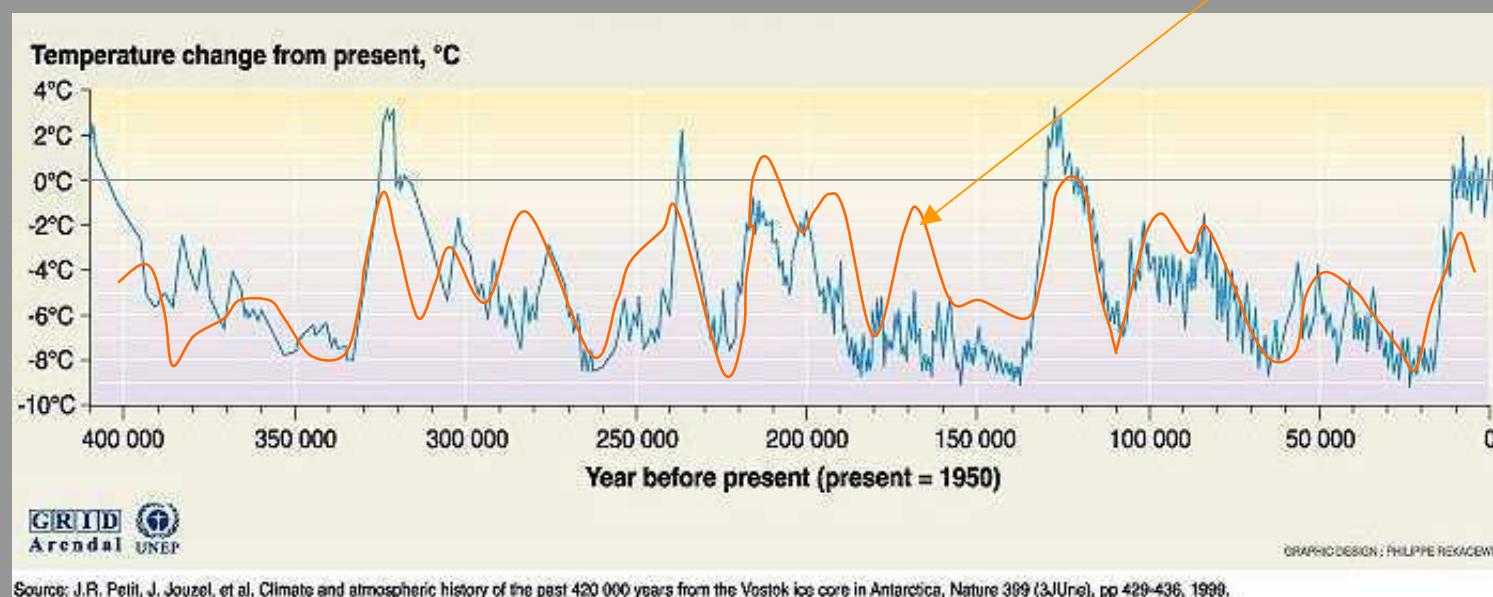
22° to 24.5°, over
41,000 years

Variations in insolation (in watts per square meter) determined from the variation in Earth's orbital elements (Barron, 1994, figure 13).

Variations in Earth's orbital elements, eccentricity, tilt (obliquity), and time of perihelion (precession of the equinoxes) computed for the last 500,000 years with a computer program written by Tamara Ledley and Starley Thompson (Barron, 1994, figure 12).



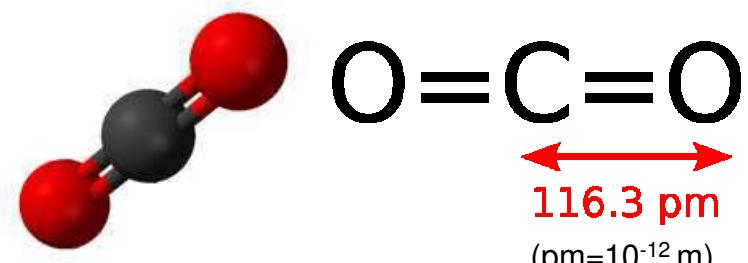
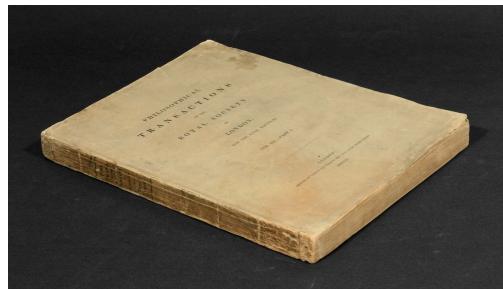
Insolation according to Milankovitch parameters





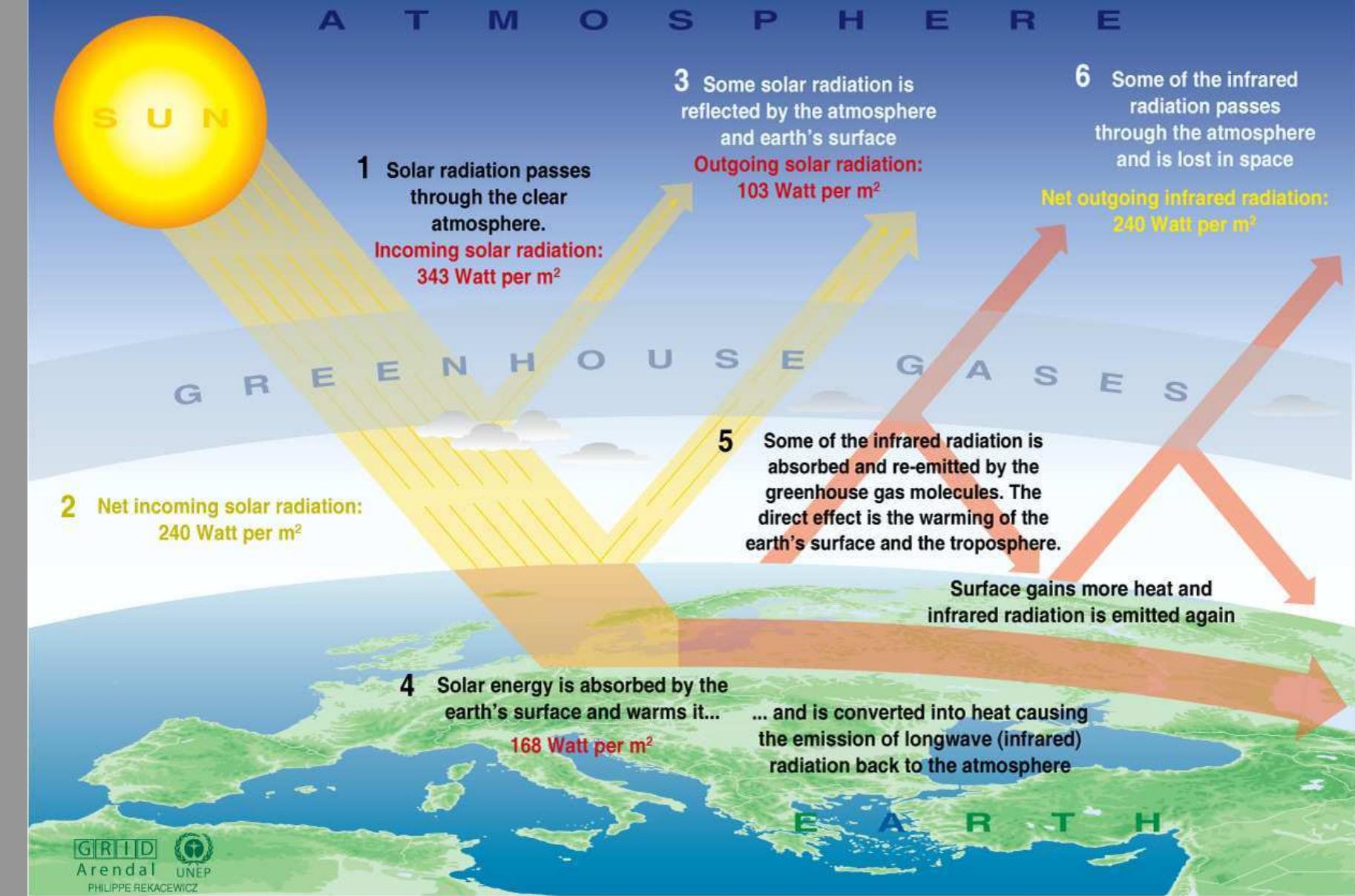
John
Tyndall
(1859)

- He established for the first time that molecules of WV, CO₂, N₂O, CH₄ y O₃ exhibit different absorption properties when IR radiation travels through them
- Changes in the amount of any of the radiatively active constituents of the atmosphere “could have produced all the mutations of climate which the researches of geologists reveal”
(1861)



Tyndall, J. (1861): On the Absorption and Radiation of Heat by Gases and Vapours, Philosophical Magazine (Ser.4) 22, pp 276-277

The Greenhouse effect

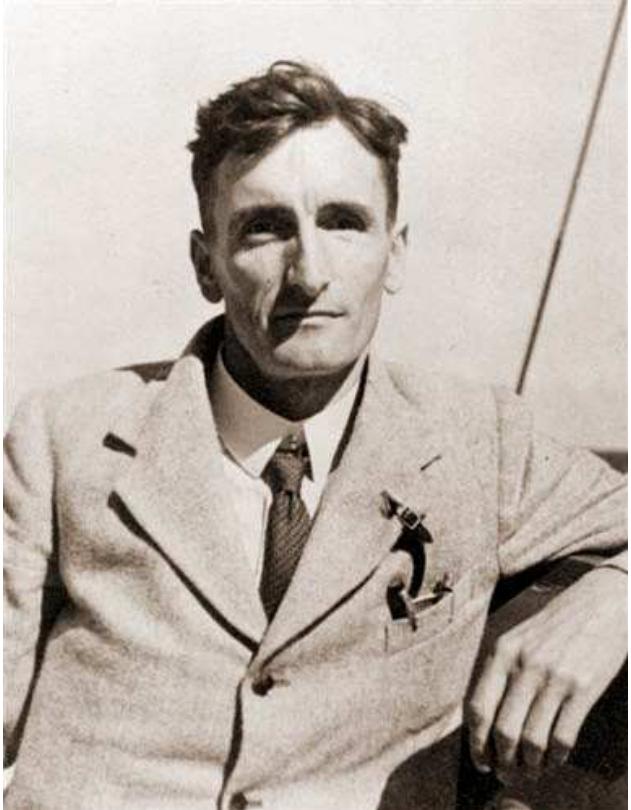


Sources: Okanagan university college in Canada, Department of geography, University of Oxford, school of geography; United States Environmental Protection Agency (EPA), Washington; Climate change 1995, The science of climate change, contribution of working group 1 to the second assessment report of the intergovernmental panel on climate change, UNEP and WMO, Cambridge university press, 1996.



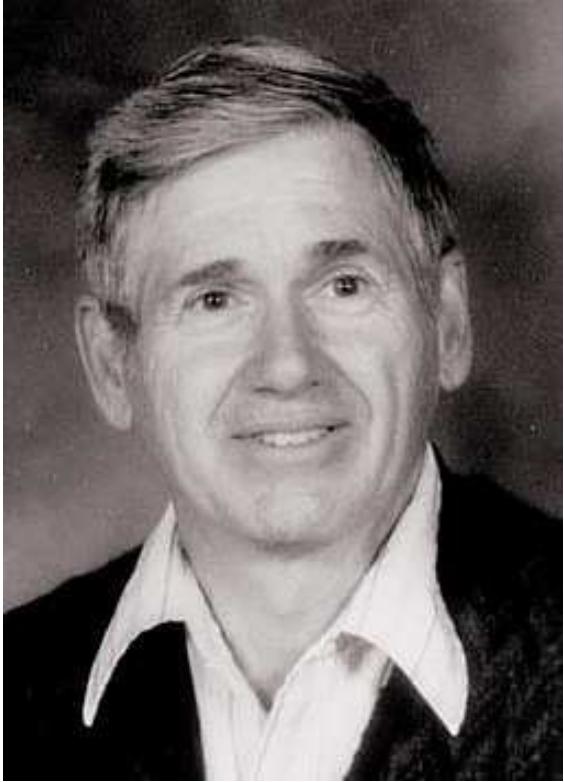
Svante Arrhenius (1896)

- He showed that doubling the concentration of CO₂ in the global atmosphere would lead to changes in average surface air temperature of between about 4°-5°C (**2°-4.5°**, IPCC 2007!!)
- Combustion of coal might induce “a noticeable increase” in atmospheric CO₂ over the course of the years



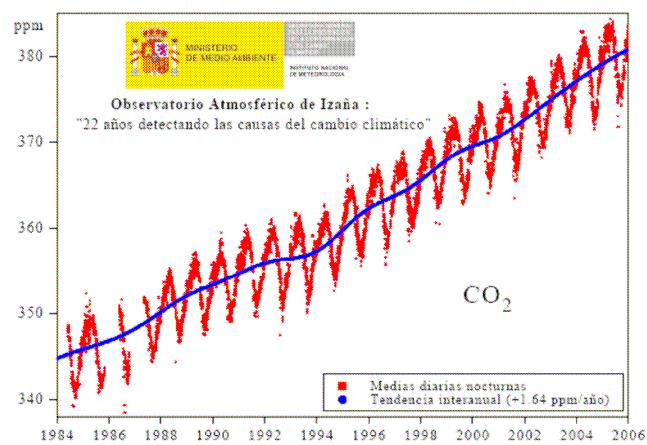
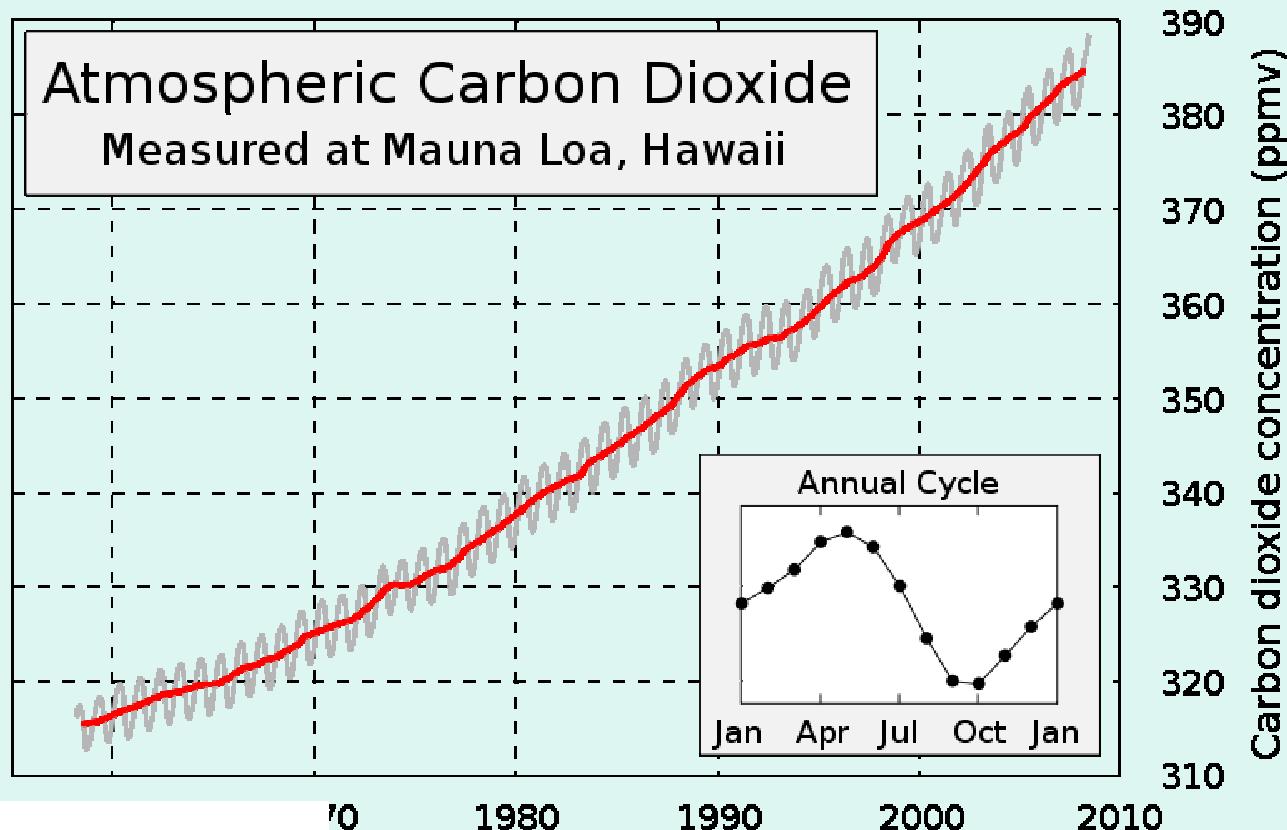
Guy S.
Callendar
(1938)

- “As man is now changing the composition of the atmosphere at a rate which must be very exceptional on the geological time-scale, it is natural to seek for the probable effects of such a change”.
- First attempt at detecting and attributing large-scale climate change to human-induced emissions of GHGs
- He linked together the 3 pillars of the idea of anthropogenic climate change: (i) CO₂ as GHG; (ii) rising concentration of CO₂; (iii) increase in world temperature



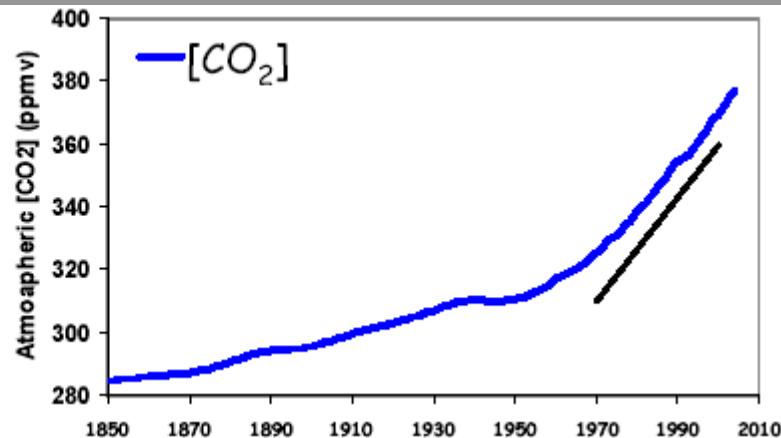
Charles D.
Keeling
(1957)

- Which is the fate of the CO₂ molecules emitted into the atmosphere?
- First incontrovertible evidence for the contemporary increase of CO₂ concentration in the atmosphere



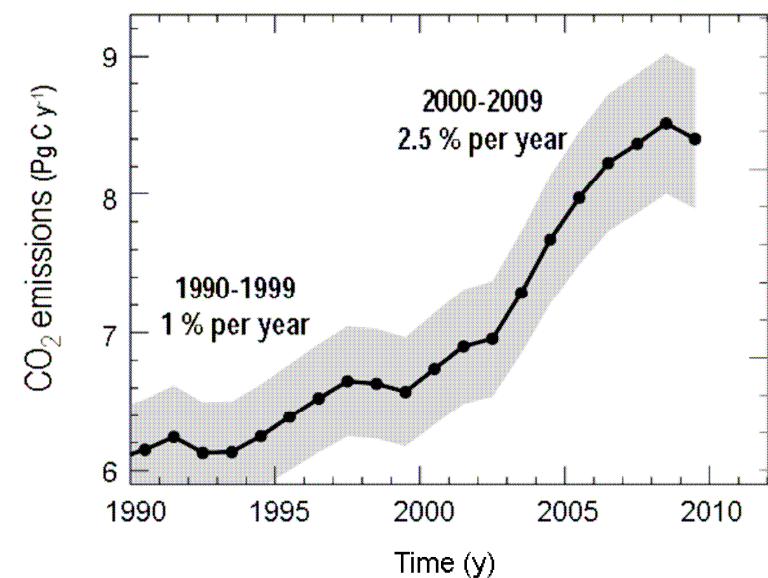
Concentración de CO₂ en la atmósfera

Year 2007
Atmospheric CO₂
concentration:
382.6 ppm
35% above pre-industrial

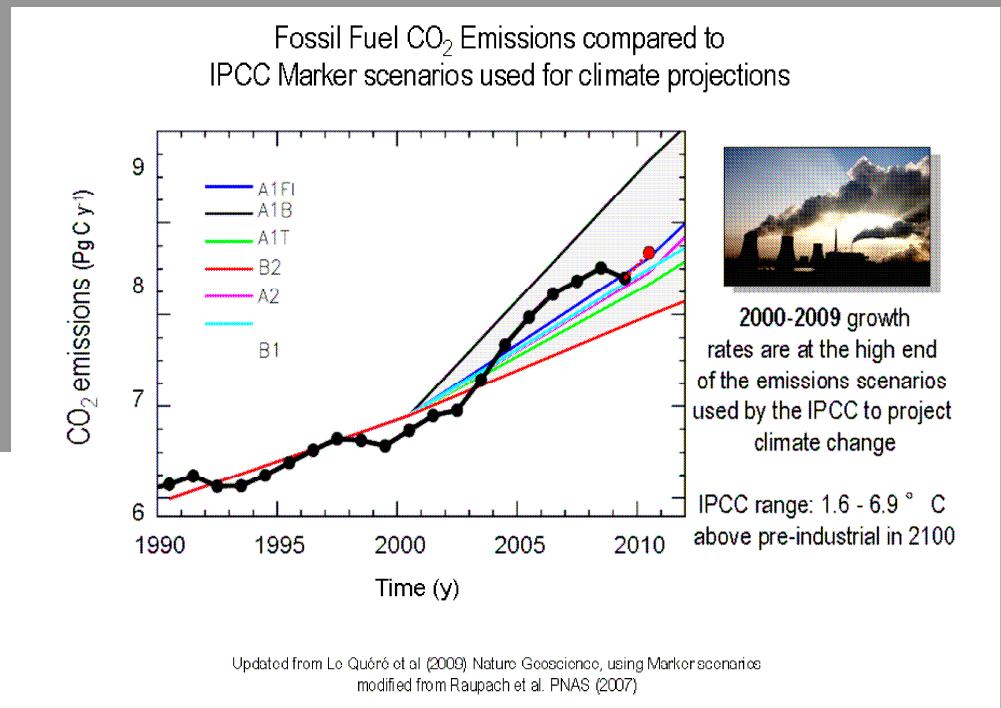


1970 - 1979: 1.3 ppm y^{-1}
1980 - 1989: 1.6 ppm y^{-1}
1990 - 1999: 1.5 ppm y^{-1}
2000 - 2006: 1.9 ppm y^{-1}

Emisiones CO₂



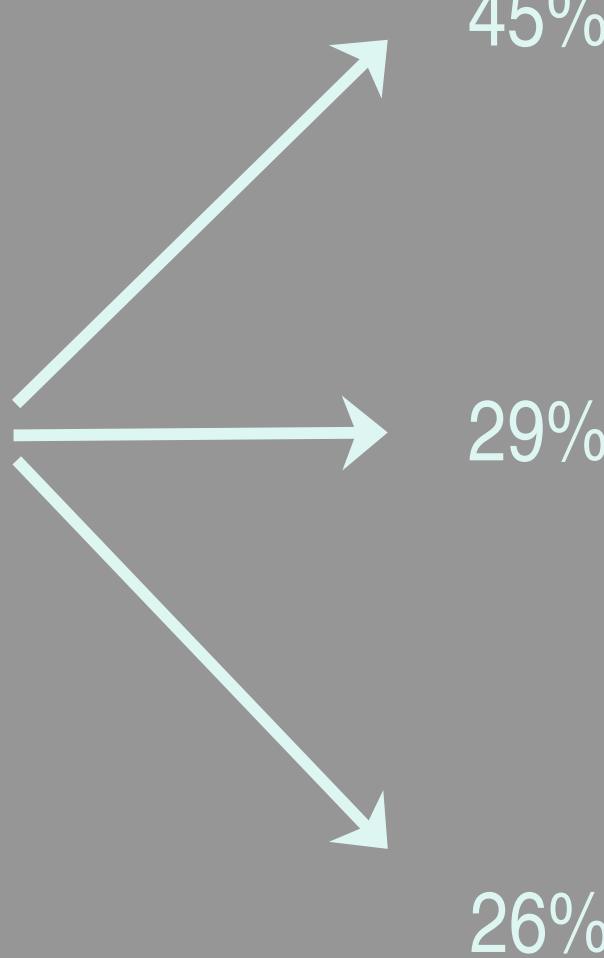
CDIAC; Friedlingstein et al. 2010, Nature Geoscience



Fate of anthropogenic CO₂ emissions (2000-2008)



+





- First model experiment which explicitly simulated the three-dimensional response of global climate to a doubling of atmospheric CO₂ concentration (2.9°C)

Syukuro
Manabe
(1975)

Modelos climáticos

$$\frac{du}{dt} = -\frac{1}{\rho a \cos \varphi} \frac{\partial p}{\partial \lambda} + fv + uv \frac{\tan \varphi}{a} + F_\lambda$$

$$\frac{dv}{dt} = \frac{1}{\rho a} \frac{\partial p}{\partial \varphi} - fu - u^2 \frac{\tan \varphi}{a} + F_\varphi$$

$$\frac{\partial p}{\partial z} = -\rho g$$

$$\frac{\partial \rho}{\partial t} = -\nabla \cdot \rho \mathbf{V}$$

$$C_p \frac{d\Theta}{dt} = \frac{\Theta}{T} Q$$

$$p = R\rho T$$



u, v, w, p, ρ, T

Ecuaciones atmósfera

Los modelos climáticos son programas informáticos basados en las ecuaciones que describen la evolución de los distintos componentes del sistema climático: atmósfera, océano, hielos, biosfera,

...

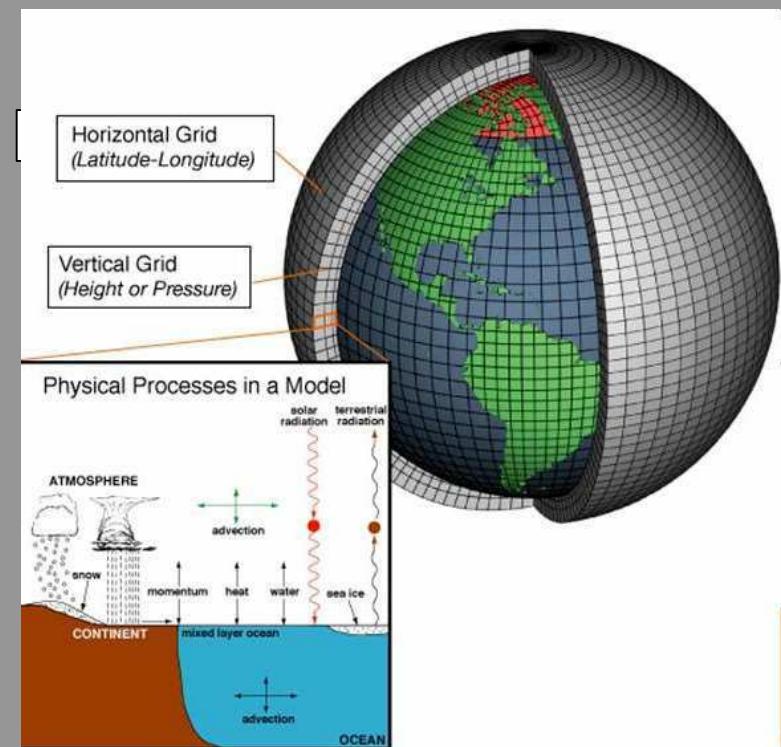
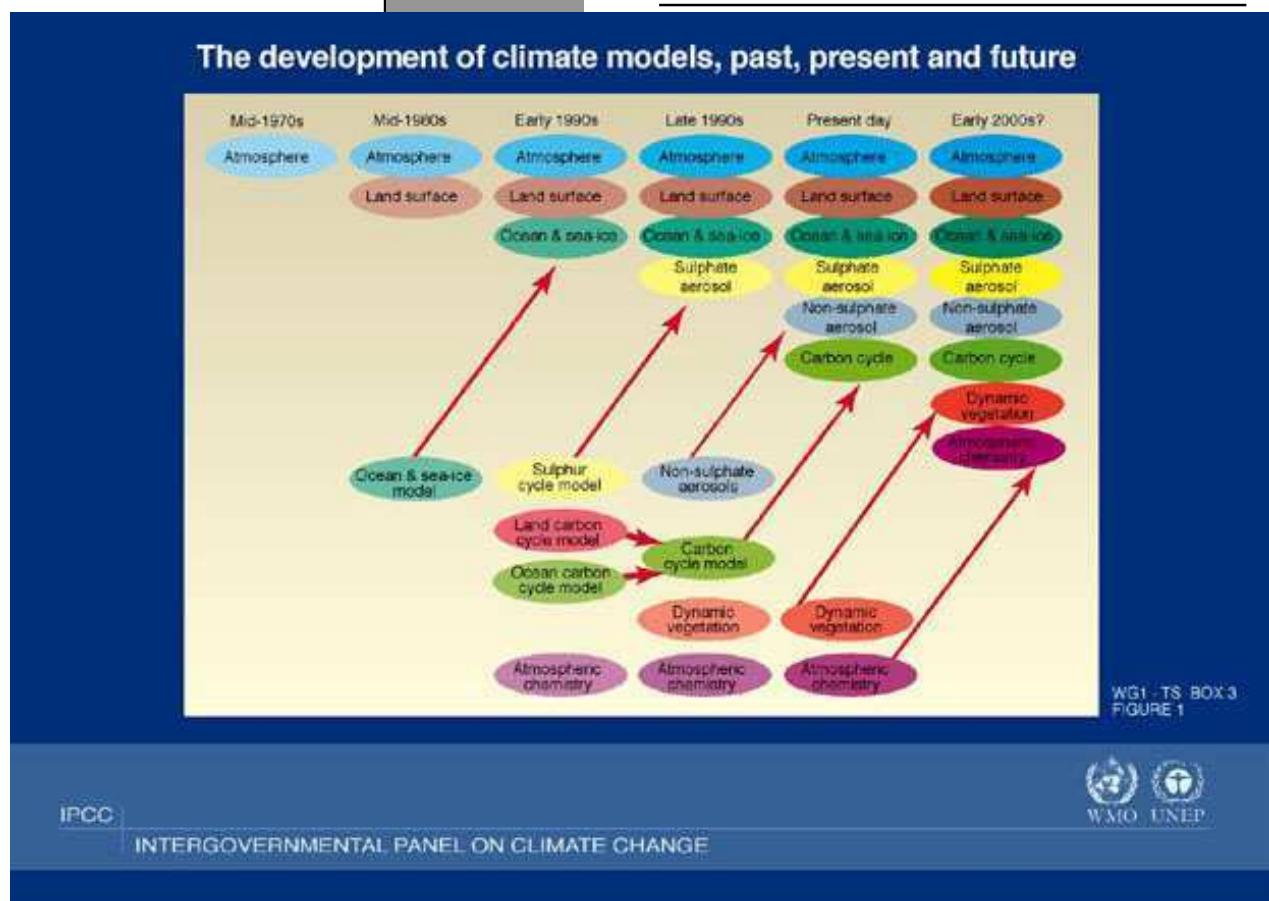
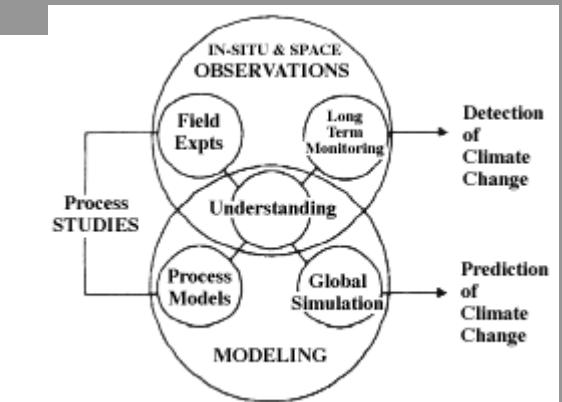
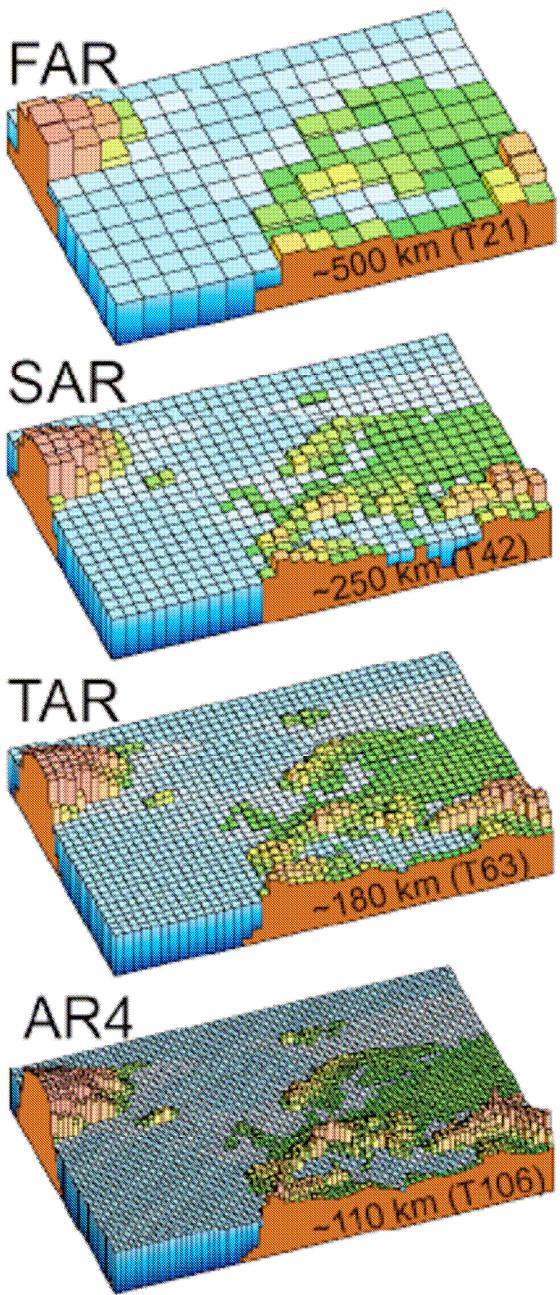


Table II. Historical evolution of climate models

Decade and landmark papers	Climate model status
≤1969 Manabe and Möller (1961) Manabe and Strickler (1964) Sellers (1969) Budyko (1969)	Numerical weather forecasts extended RC models developed Dynamics and radiation virtually separate EBMs newly described
1969–1981 Manabe and Bryan (1969) Green (1970), Stone (1973) Manabe and Wetherald (1975) CLIMAP (1981)	Multi-layer oceans added to GCMs SD models developed Greenhouse modelling with GCMs Palaeo datasets first employed for 'validation'
1981–1989 Hansen <i>et al.</i> (1981) Sellers <i>et al.</i> (1986) Oort and Peixoto (1983) Luther <i>et al.</i> (1988)	GCMs becoming predominant model type Surge in computational power and capacity Satellites generate global observations Model intercomparisons suggested
1989–1999 Houghton <i>et al.</i> (1990) Semtner and Chervin (1992) Flato and Hibler (1992) Cubasch <i>et al.</i> (1994) Santer <i>et al.</i> (1996)	Simpler models required by IPCC OAGCMs established but need flux correction Sea-ice and land-surface components evolving First ocean-atmosphere coupled ensemble Validation and attribution first described
2000s ???	EMICs as important as GCMs Past climate simulations re-emerging for testing Observational need driven by evaluation demand Policy needs a major driver of numerical models



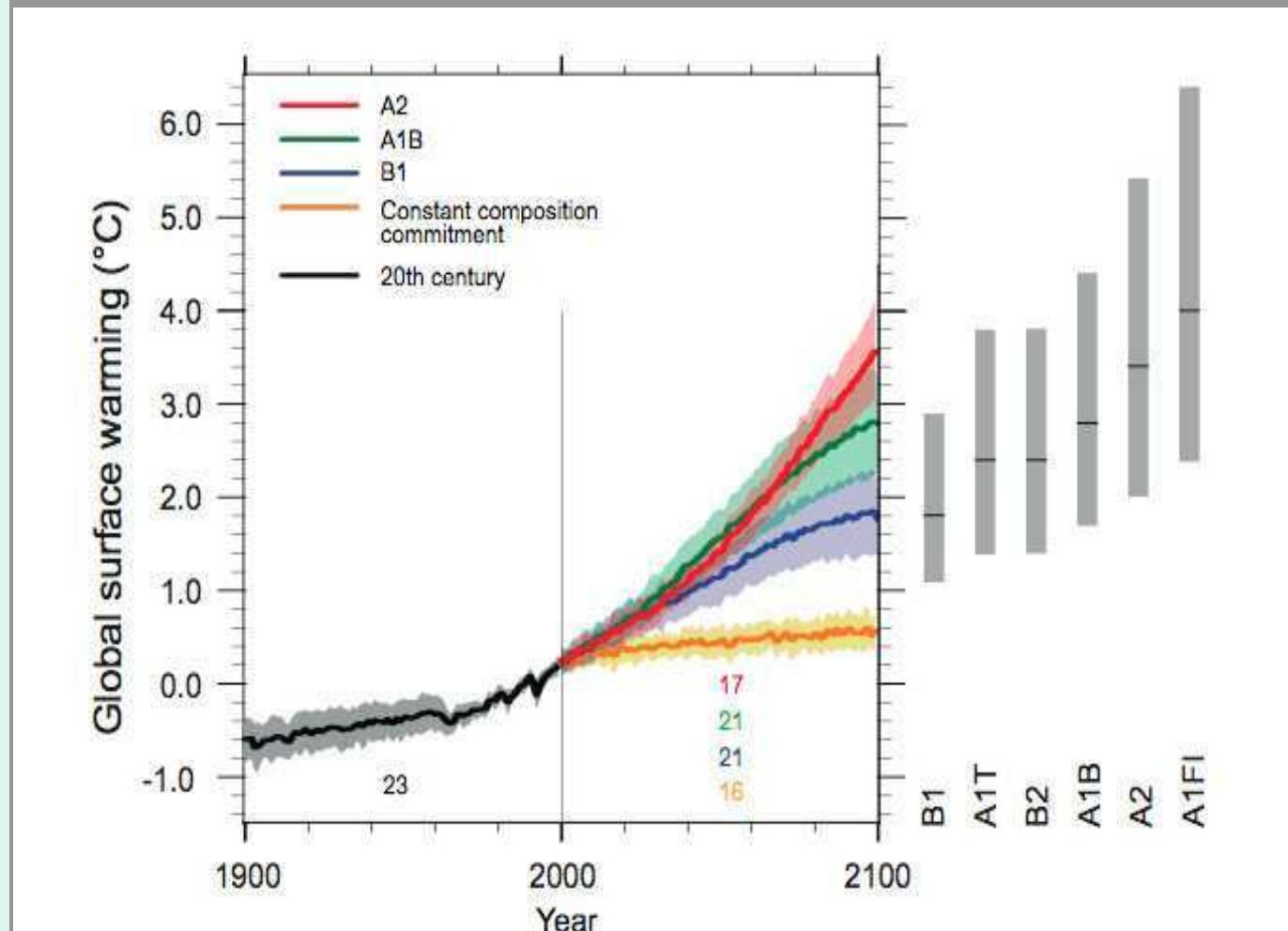
Proyecciones de cambios futuros en el clima

(escenarios sin compromisos políticos!!!)

* Mejor estimación para escenario bajo (B1) es 1.8 °C (rango probable 1.1-2.9 °C), y para escenario alto (A1FI) es 4.0 °C (rango probable 2.4-6.4 °C).

* Generalmente consistente con el rango citado para SRES en TAR pero no directamente comparable

* Dos próximas décadas aprox. 0.2°/decada para muchos de los SRES



(AR4, 2007)

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